

# Skin Surface Temperatures

First Look at analysis of COSMICS  
data

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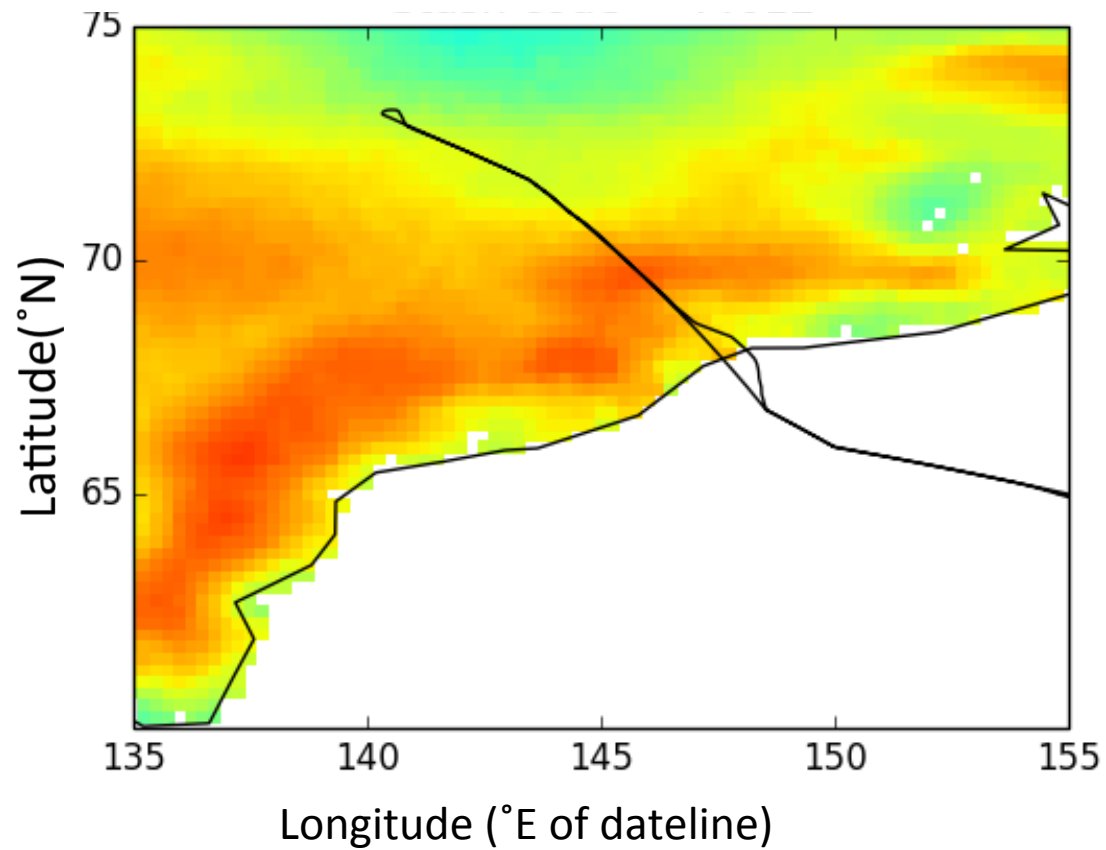
# Motivation

- Past work from arctic campaigns has focused on MW surface emission
- This analysis focuses on skin temperature retrievals to facilitate FIR retrievals of emissivity using TAFTS.
- Some MW retrievals are presented but this MW analysis is far from complete.

# Three ways to estimate surface temperature

- (1) Onboard Heimann estimate of surface temperature
  - IR brightness temperature (8-14  $\mu\text{m}$ )
  - High temporal resolution (1 or 64 Hz)
  - Must be calibrated to get surface temperature
- (2) ARIES IR Interferometer
  - $T_B$  500 to 3500  $\text{cm}^{-1}$  at 1  $\text{cm}^{-1}$
  - Nadir and zenith and calibration measurements every X minutes
  - Simultaneous retrievals of surface temperature and emissivity
- MARSS:  $T_{\text{eff}}$  estimated from 183 GHz channels
  - Each MARSS footprint (nadir every 4 sec)
  - Only uses MARSS data to derive  $T_{\text{eff}}$

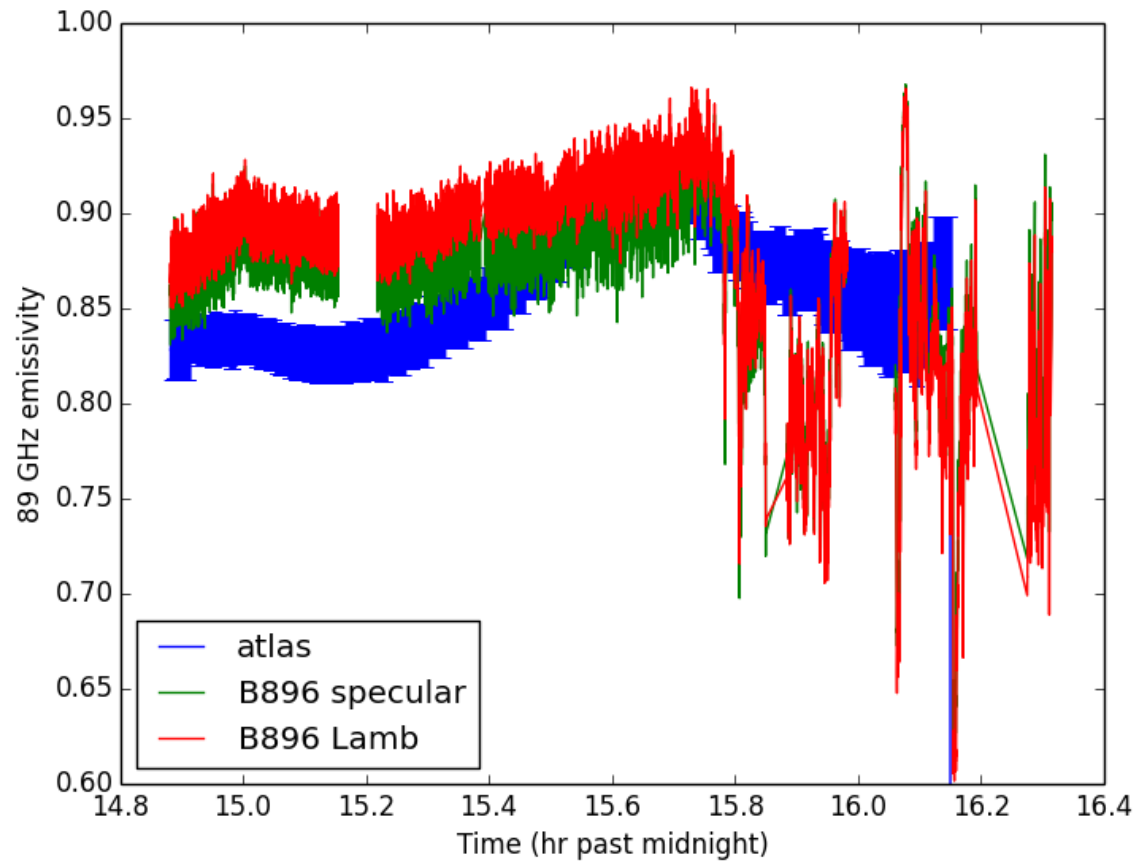
# B896: First sortie over Greenland



0.60 0.65 0.70 0.75 0.80 0.85 0.90 0.95

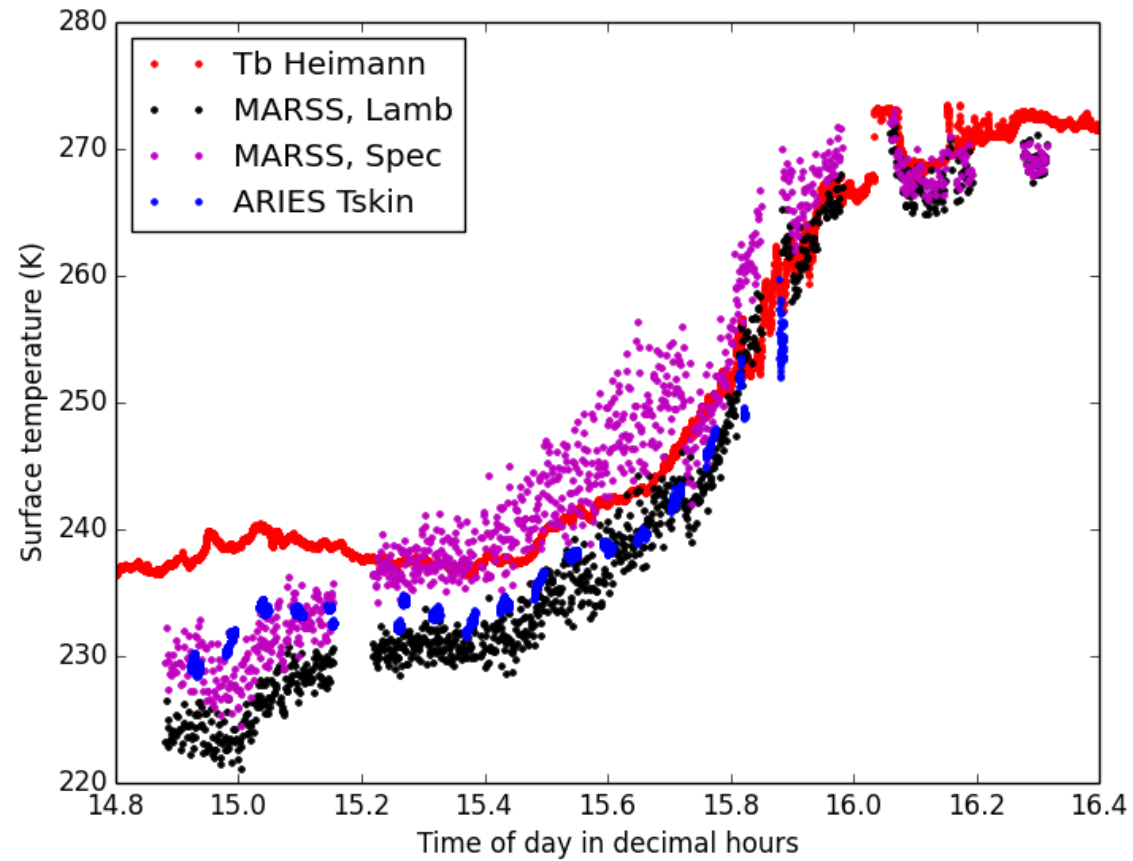
89 GHz Emissivity (Karbou atlas)

# B896: Comparison of 89 GHz emissivity

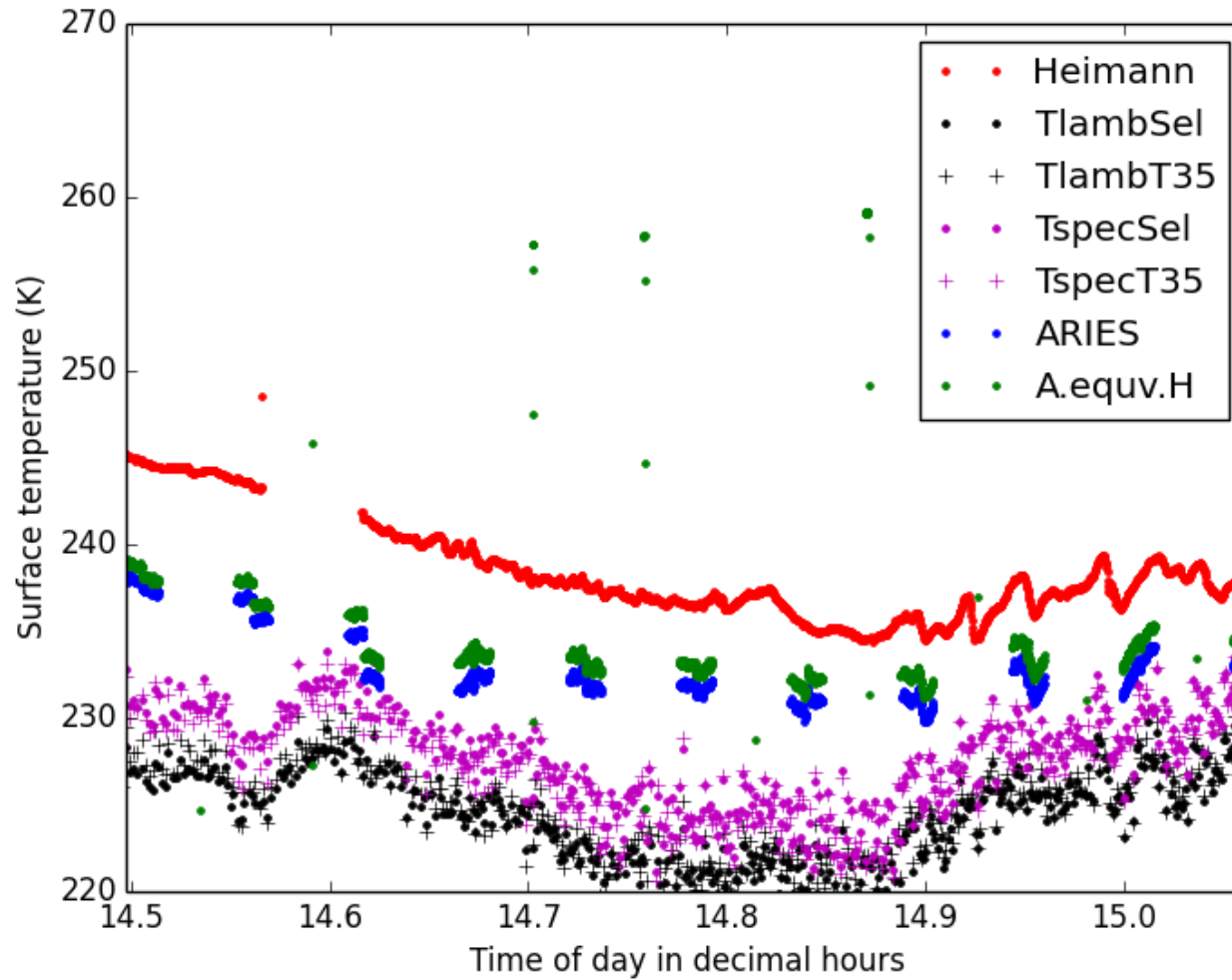


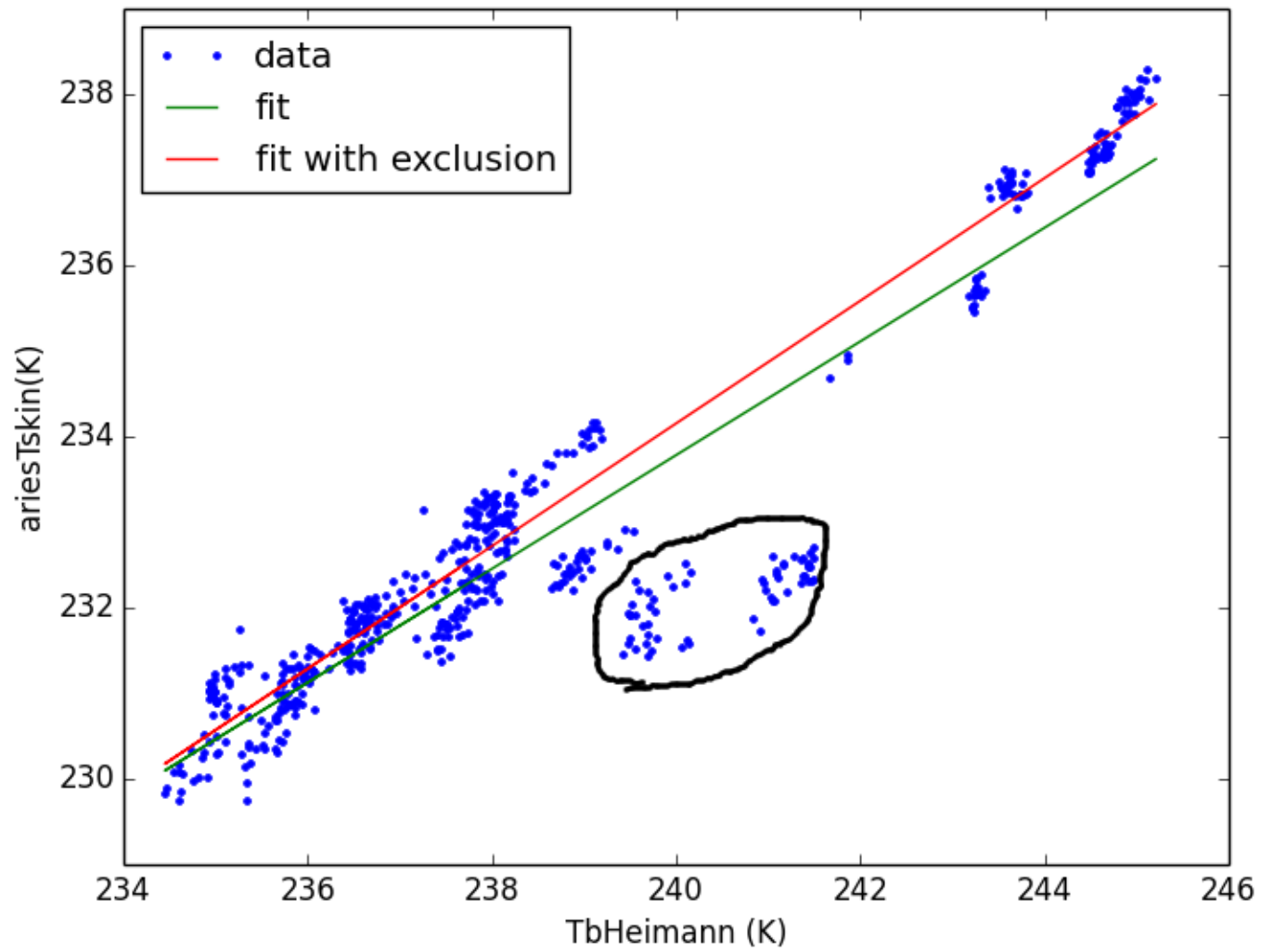
Karbou atlas in Blue

# B896: comparison of surface temps



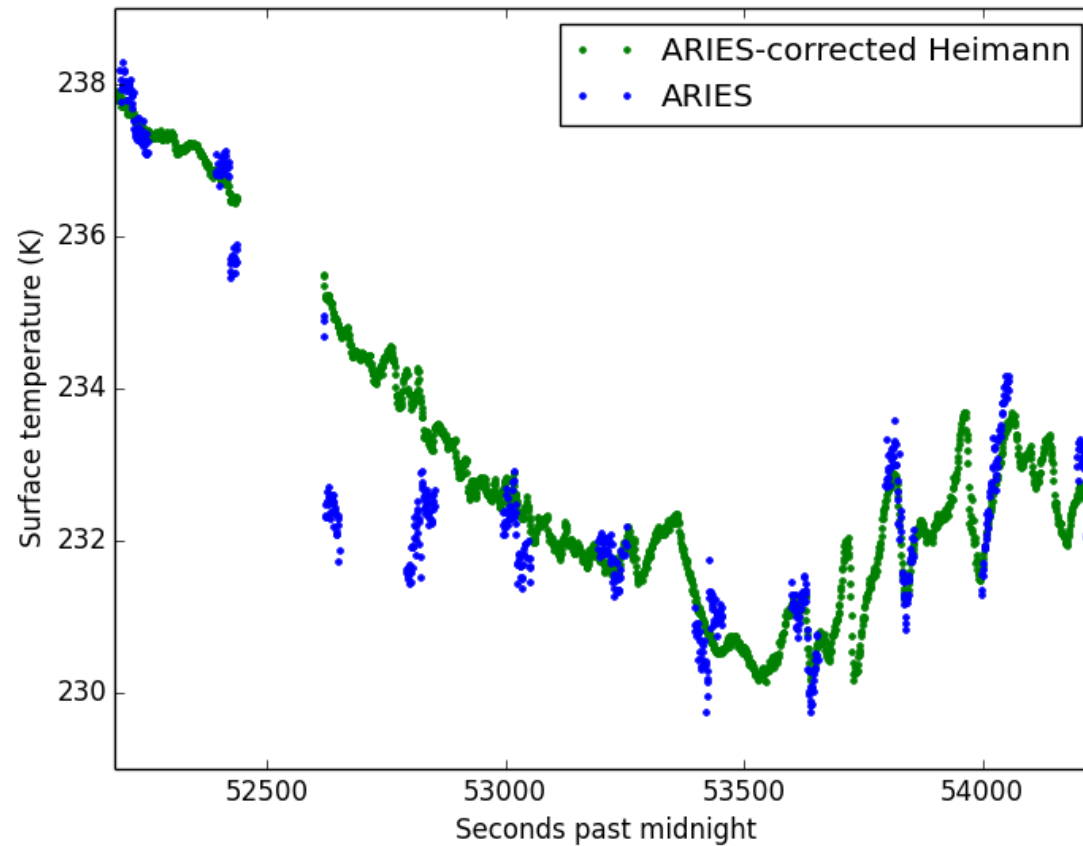
# B898 Surface Temperatures







# ARIES-corrected Heimann



This surface temp used in retrievals of TAFTS emissivities to be presented at a later time.

Questions or comments

# Determining emissivity and effective temperature

- Need
  - Measurements of  $T_{Bn}$  and  $T_{Bz}$  on 183 GHz sounding channels ( $183\pm 1$ ,  $183\pm 3$  and  $183\pm 7$  GHz).
  - Measurements of temperature and water vapor profile between the platform and surface.
- Assume linear emissivity gradient between 175 and 191 GHz
  - $e(183\pm 7) \equiv e(183\pm 1) \equiv e(183\pm 3) \equiv e(183 \text{ GHz})$
- Use simple clear skies radiative transfer to extrapolate measurements at height to the surface (small effect in dry air)

# Selbach: 183 GHz effective temperature and emissivity

- Uses all three 183 GHz channels.
- Simple clear skies radiative transfer model

$$T_{Bn} = e_s T_{eff} \exp(-\tau) - (1 - e_s) T_d \exp(-\tau) + T_a$$

$$T_d = T_{Bz} \exp(-\tau) + T_a$$

$$T_a = (1 - \exp(-\tau)) T_m$$

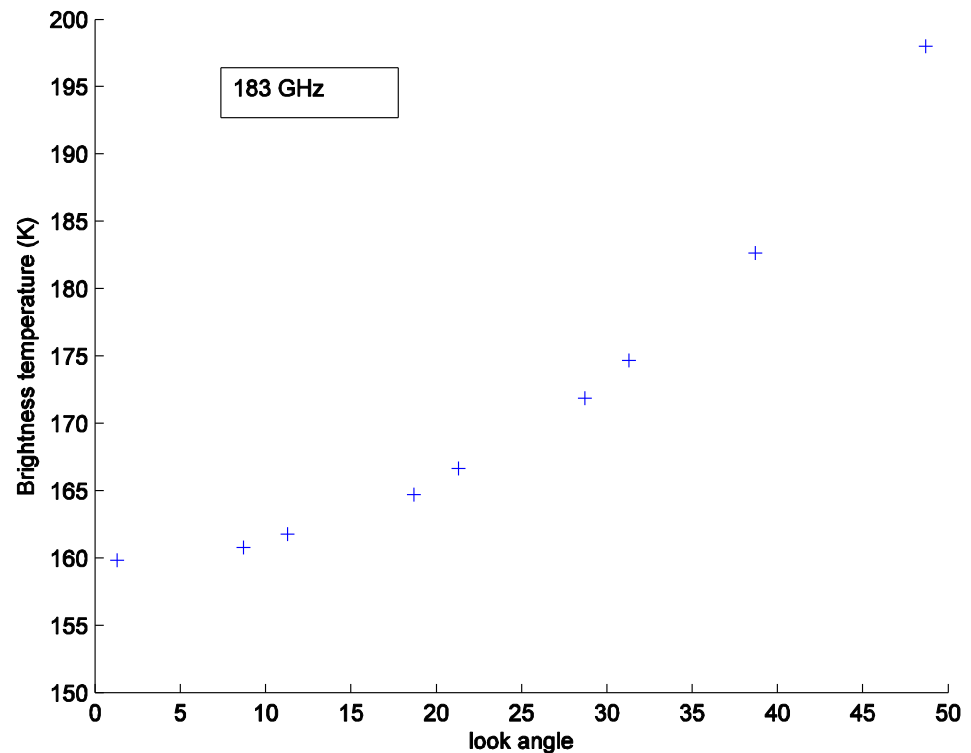
- $e_s$  and  $T_{eff}$  -- surface emissivity and effective temperature.
- $T_{Bzi}$  and  $T_{Bni}$  -- measured zenith and nadir viewing brightness temperatures in channel  $i$ .
- $T_m$  -- mean atmospheric temperature under the aircraft.
- $\tau_i$  is the opacity in channel  $i$ . Determined with ARTS using dropsonde profiles.
- Differences between modelled and observed  $T_{Bn}$ 's on the three 183 GHz channels are analytically minimized in cost function.
  - Closed form solution:  $T_{eff}$  and  $e_s$  at 183 GHz

# Assumptions about reflection

- Up to this point all results presented use specular reflection assumption
- Recent literature (Mätzler, 2005; Mätzler and Rosenkranz, 2007) over snow covered surfaces
  - specularity not a good assumption for near-nadir viewing satellite instruments such as AMSU-A and AMSU-B.
  - Reflection more diffuse in character.
- MARSS scans between  $0^\circ$  and  $50^\circ$  incidence in the upward and downward directions.
  - Near-nadir views overlap with views of other radiometers ARIES and Heimann and the radar altimeter.
  - Diffuse surface scattering characteristic important for retrieving near-nadir emissivities.
- Now demonstrate effect of diffuse surface scattering.

# MARSS Tip Curve

- Optically thin channels have  $T_{bd}$  that increases with incidence angle.
- $T_{MR}$  is the mean atmospheric temperature weighted by the absorption in each layer.
- $T_{CMB}$  is the cosmic background radiation.
- At high optical depth,  $\tau$ ,  
 $T_{bd}(\theta) = T_{MR}$



$$T_{bd}(\theta) = T_{MR} (1 - \exp(-\tau / \cos \theta)) + T_{CMB} \exp(-\tau / \cos \theta)$$

# Calculating surface scattering contribution with MARSS measurements.

$$T_d(\mu_0, \phi_0) = \frac{1}{4\pi\mu_0} \int_{2\pi} S(\mu_0, \phi_0, \mu, \phi) T_{bd}(\mu, \phi) d\Omega$$

- MARSS makes six angular measurements of  $T_{bd}$  at  $1^\circ$  to  $49^\circ$  wrt vertical in upward directions.
- Must estimate  $T_d(\mu_0, \phi_0)$  with the limited views provided by MARSS.
- There will be a contribution to  $T_d(\mu_0, \phi_0)$  from outside the scan.
  - Estimated by calculating above over theoretical ‘Tip Curve’ to estimate the proportion of integral outside of MARSS views.
  - $T_d(\mu_0, \phi_0)$  is then calculated for the MARSS measurements and corrected for partial coverage of sky.
  - These theoretical ‘Tip Curves’ are only valid for homogeneous, clear skies cases.

# Technote 35: 183 GHz effective temperature and emissivity

- Uses classical definition of emissivity ((2) below) (1)

$$T_{eff} = \frac{T_u(183 \pm 1GHz) - T_d(183 \pm 1GHz)}{e_s(183GHz)} + T_d(183 \pm 1GHz) \quad (6) \quad (1) \quad (2)$$

$$e_s(183GHz) = \frac{T_u(183 \pm 7GHz) - T_d(183 \pm 7GHz)}{T_{eff} - T_d(183 \pm 7GHz)} \quad (7) \quad (2)$$

- (1) and (2) combine to form (3)

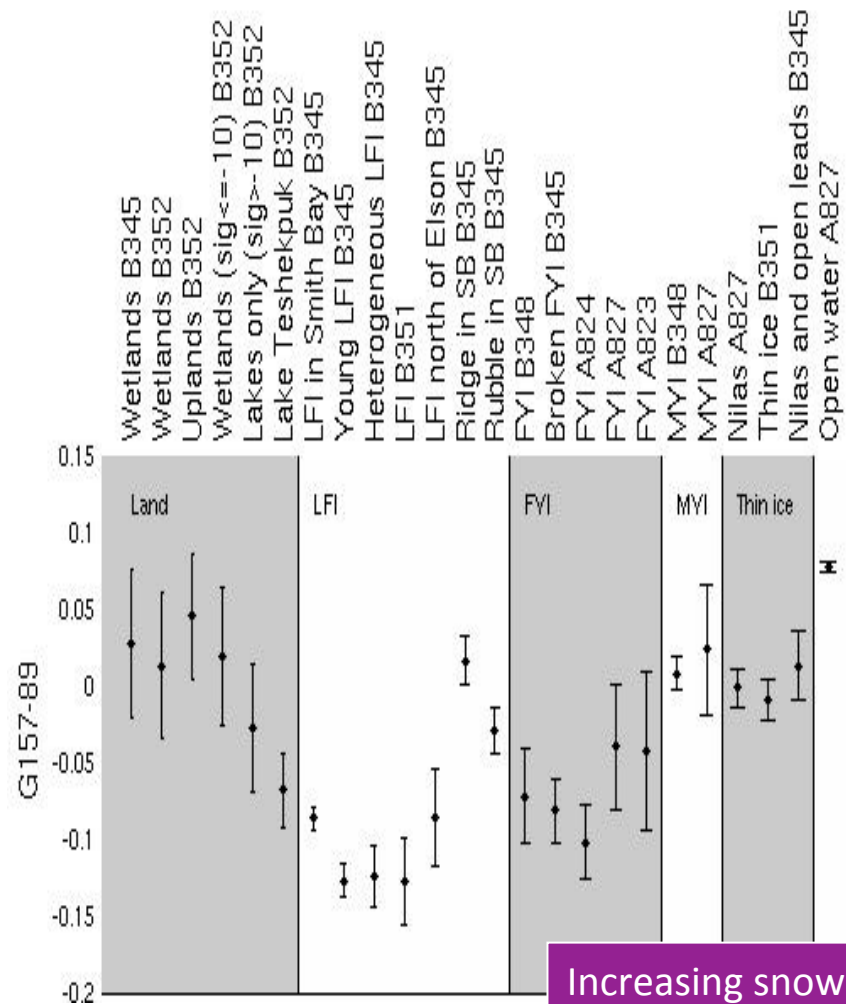
$$e_s(183) = \frac{T_u(183 \pm 7) - T_u(183 \pm 1) - T_d(183 \pm 7) + T_d(183 \pm 1)}{T_d(183 \pm 1) - T_d(183 \pm 7)}$$

- Solution of (3) used in (1) to find  $T_{eff}$
- Only uses 183±1 and 183±7 GHz channels



# Sea ice mm-wave emissivities can be related to sea ice type

- Broad surface classes in alternating grey and white
- mm-wave Lambertian emis can be related to surface type (Harlow, 2010)
  - 157–89 GHz emis difference sensitive to snow depth and stratigraphy.
  - 89 GHz radiation penetrates 10 – 20 cm in pack.
  - Penetration at 157 GHz shallower.
  - Insensitive to ice surface roughness and salinity
- Explore with MEMLS later



Increasing snow depth and depth hoar fraction