

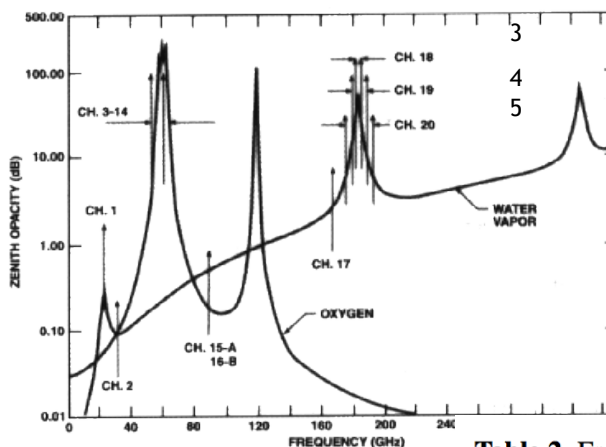
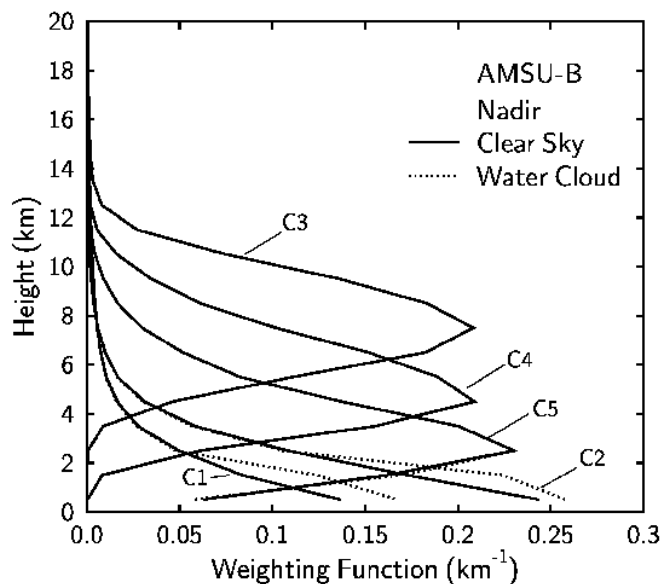
# Utilisation des observations micro-ondes passives AMSU-B/MHS pour l'étude des nuages précipitants

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## AMSU-B/MHS



**Table 2.** Frequencies of AMSU-B and MHS channels.

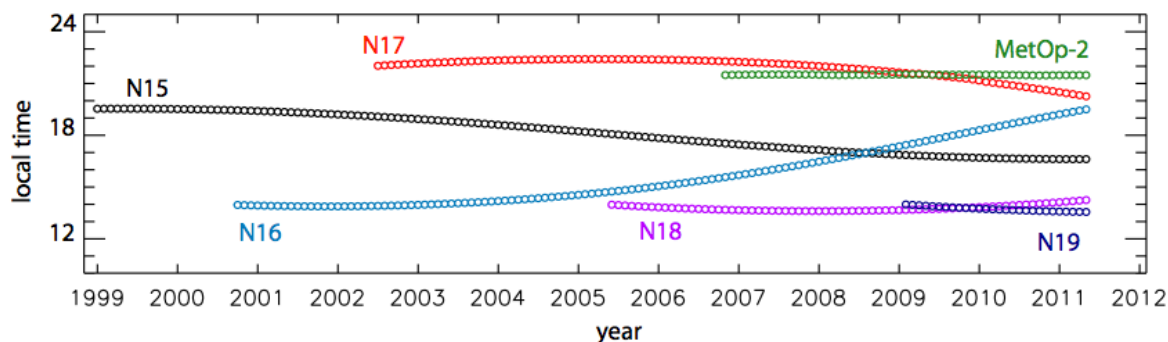
Channel Number	AMSU-B Center Freq. (GHz)	MHS Center Freq. (GHz)
1	89	89
2	150	157
3	183 ± 1	183 ± 1
4	183 ± 3	183 ± 3
5	183 ± 7	190

Platform	Radiometer	Launch Date	LTAN
NOAA-15	AMSU-B	13/05/1998	19:20
NOAA-16	AMSU-B	21/09/2000	14:24
NOAA-17	AMSU-B	24/06/2002	22:25
NOAA-18	MHS	20/05/2005	13:41
Metop-A	MHS	19/10/2006	21:30
NOAA-19	MHS	06/02/2009	13:43
NPP	ATMS	28/10/2011	13 :30
Metop-B	MHS	17/10/2012	21:30

**Tableau 1:** Liste des plateformes ayant à leur bord les radiomètres AMSU-B/MHS et caractéristiques principales : LTAN (Local Time Ascending Node- Heure locale du nœud ascendant) au moment du lancement et en décembre 2010 (dû à des dérives de l'orbite).

## WHY AMSU-B/MHS?

- AMSU-B water vapor channels have unique advantages, such as:
  - high sensitivity to frozen hydro-meteors in precipitating clouds [e.g., Burns et al., 1997; Wang et al., 1997; Bennartz and Bauer, 2003],
  - different sensitivities to vertical distributions of frozen hydrometeors [e.g., Burns et al., 1997; Bennartz and Bauer, 2003],
  - and negligible impact of the ground surface and liquid water cloud in the lower to middle troposphere [e.g., Muller et al., 1994; Wang et al., 1997; Greenwald and Christopher, 2002; Bennartz and Bauer, 2003].
- AMSU has improved horizontal and spectral resolutions compared to its predecessor (MSU). Resolution spatiale: 16 km, fauchée 2200 km.
- Continuous coverage: AMSU-B/MHS instrument has been collecting data onboard NOAA and METOP satellites since 1999.



## STRATEGY

There is no attempt to determine a rain rate as proposed by other authors (e.g. Kongoli et al., 2007; Vila et al., 2007; Surussawadee and Staelin, 2009; Di Tomaso et al., 2009; Laviola and Levizzani, 2011).

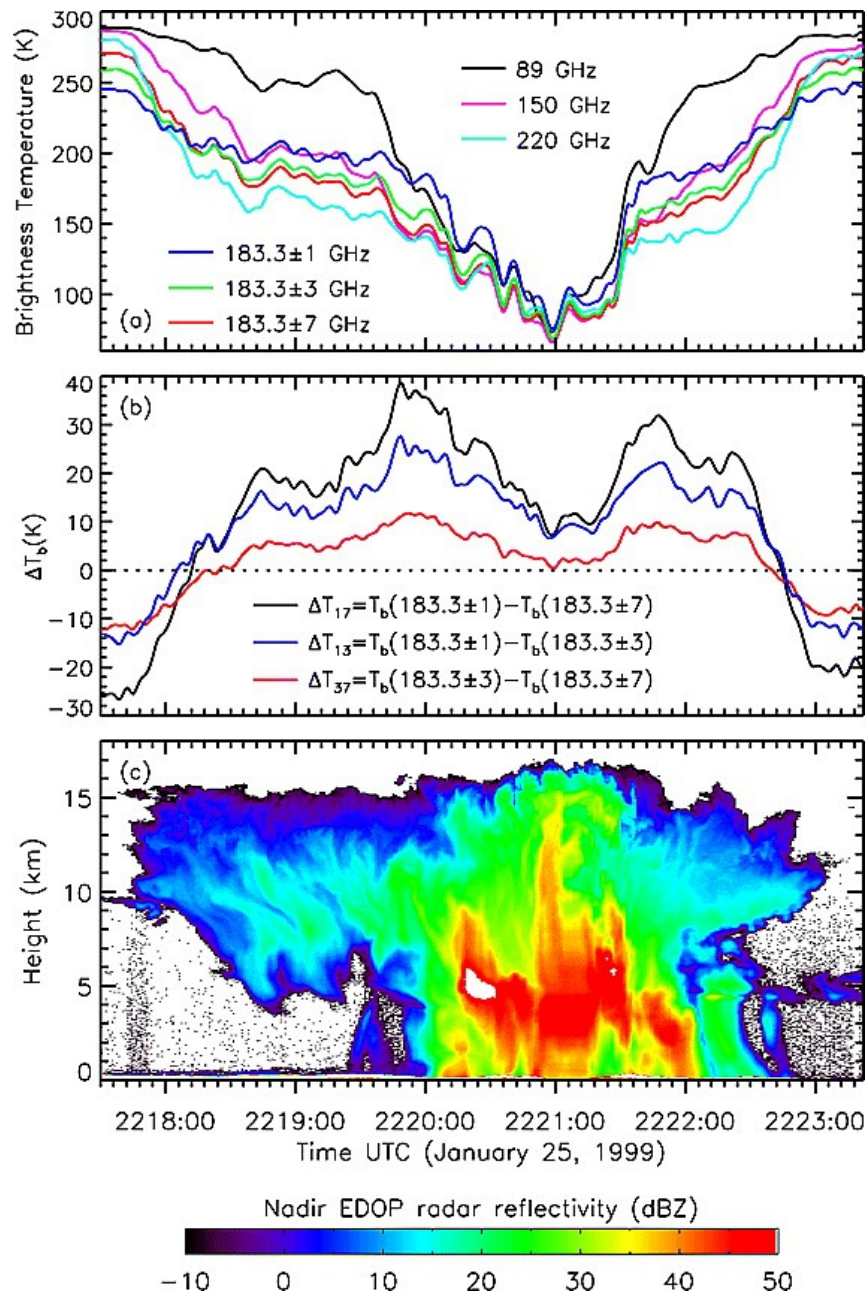
Two reasons:

- the **weak correlation** of **scattering intensity** and **instantaneous rain rate** at the surface (e.g. Bennartz and Petty, 2001).
- **saturation issues** (e.g. Lima et al., 2007) might indeed lead to an under-estimation of large rain rates associated with severe events, which are of special interest in the HYMEX project.

We have favoured an approach in which we detect precipitating clouds / deep convection occurrences.

In the following, we show:

- 3 diagnostics: MR (“Moderate rain”), DC (“Deep convection”), COV (“Convective Overshootings”)
- Some applications of these diagnostics.



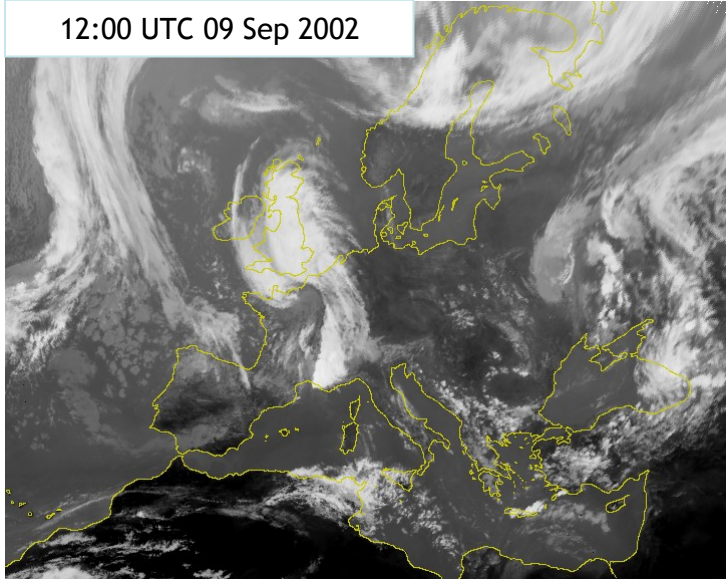
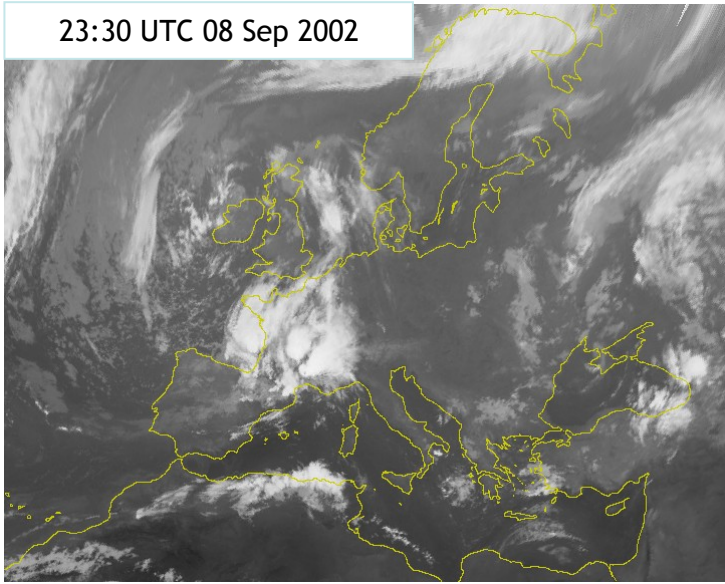
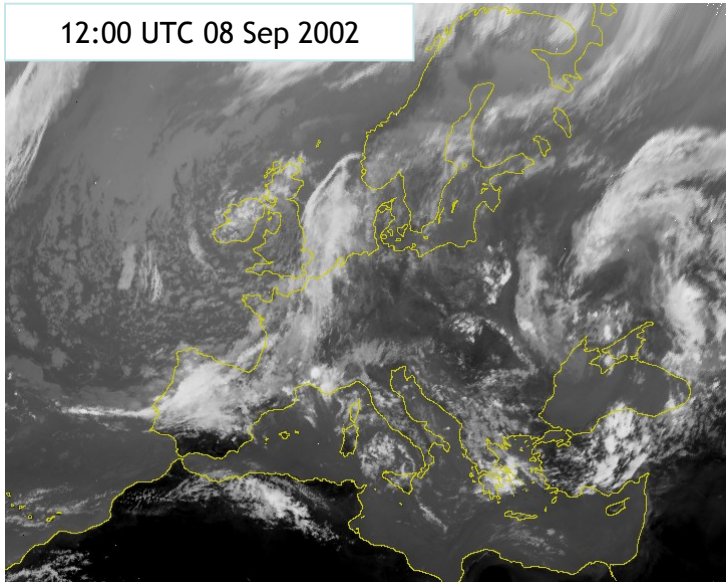
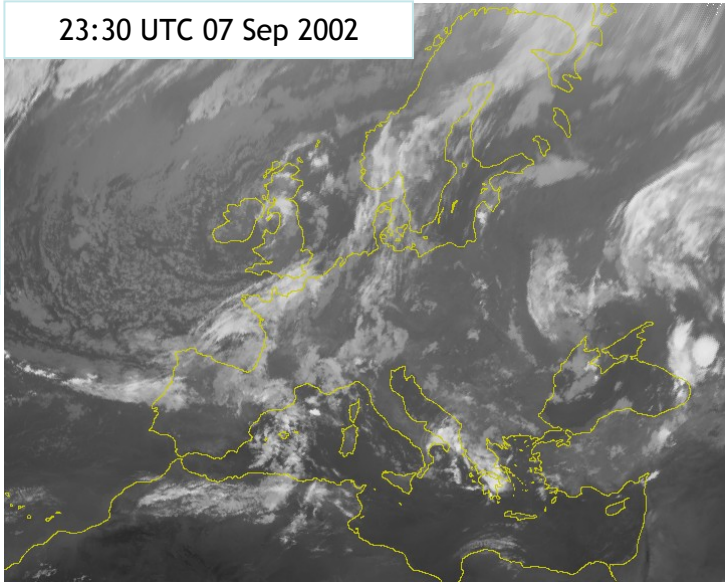
2 critères différents pour:  
 -Deep convection (DC)  
 $B3m5, B3m4, B4m5, \text{all} > \text{or} = 0$

-Convective Overshootings (COV)  
 $B3m5 > B3m4 > B4m5 > \text{or} = 0$

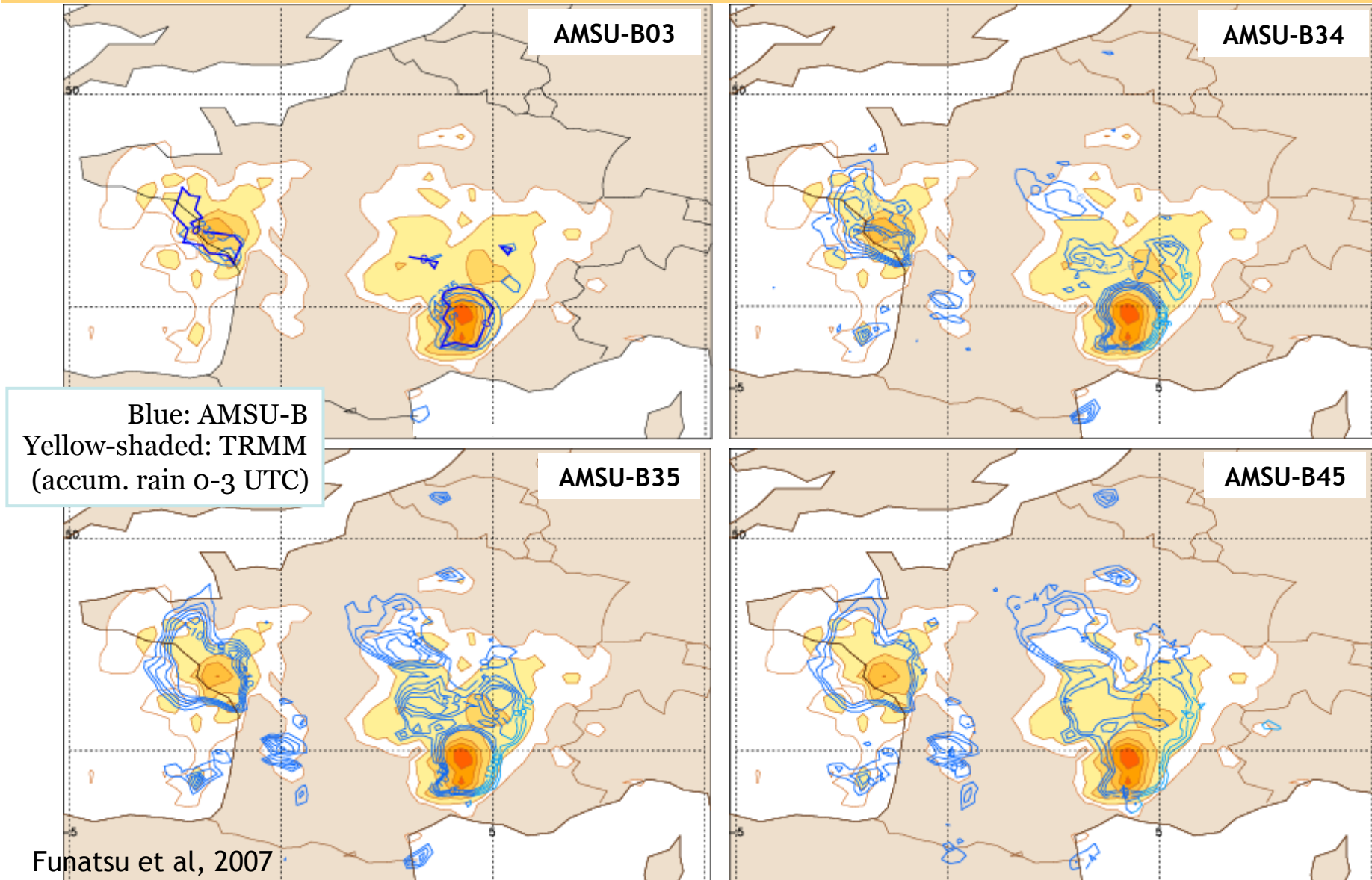
Hong et al., 2005  
 Detection of tropical deep  
 convective clouds from AMSU-B  
 vapour channels measurements

# Case Study: Nîmes-Marseille ( 8-9 September 2002)

Meteosat  
IR



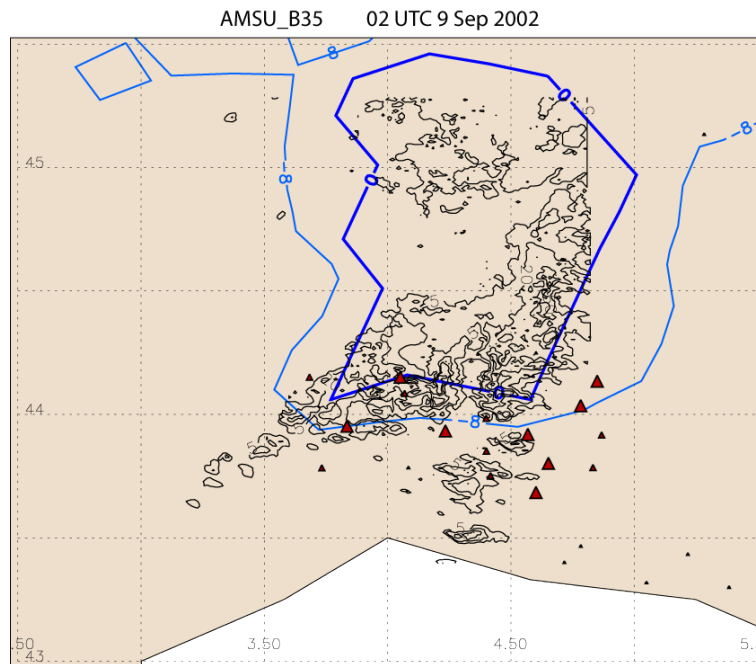
# Precipitation and AMSU-B: Comparison with TRMM<sub>3B42</sub> Nîmes-Marseille – 02 UTC 9 September 2002



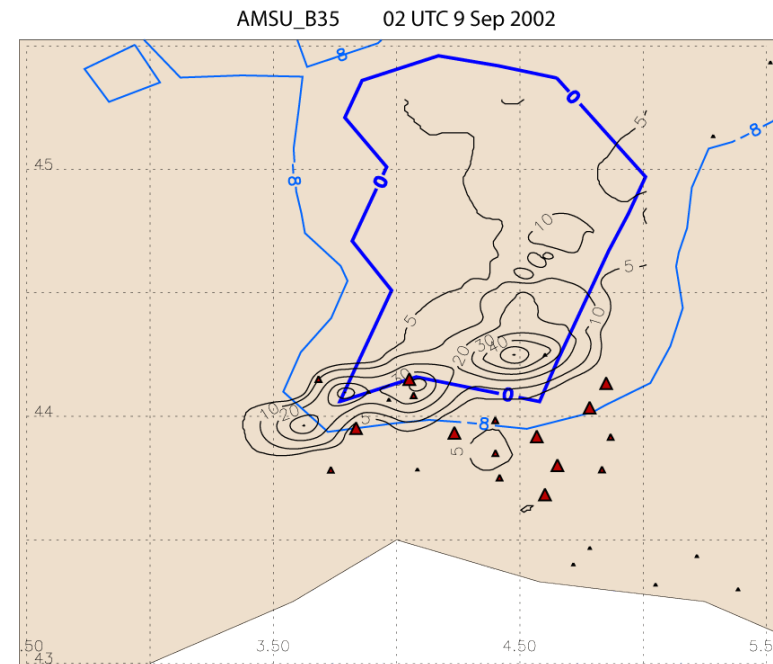
# Precipitation and AMSU-B: Comparison with Ground Data

## Nîmes-Marseille – 02UTC 9 Septembre 2002

### AMSU-B [3-5] and DCT



Thin black – radar (precip intensity in mm/h at 02:30UTC)



Black – accum. precip between 2-3 UTC

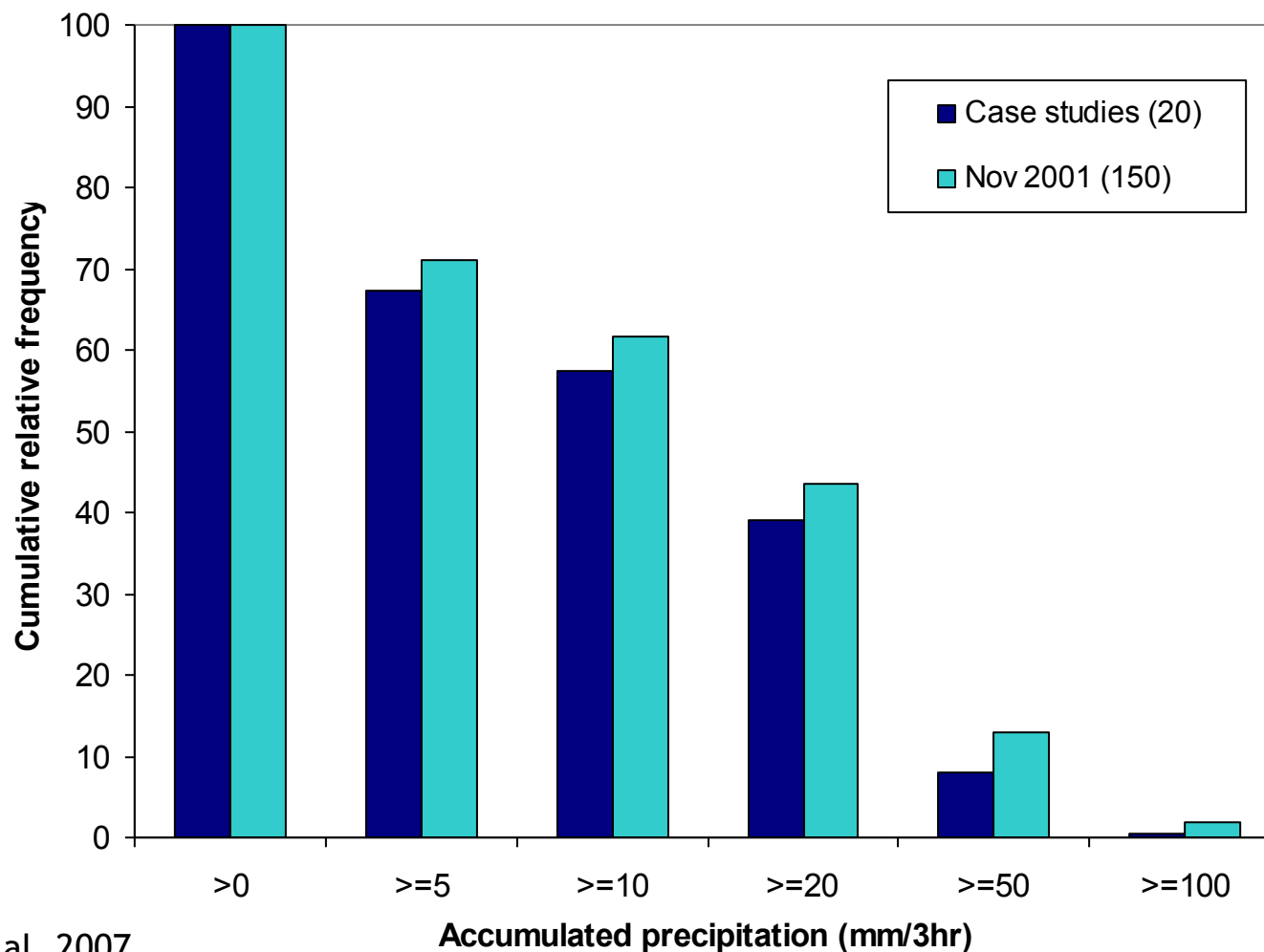
▲ - station SQR (precip. accum. during 24h)

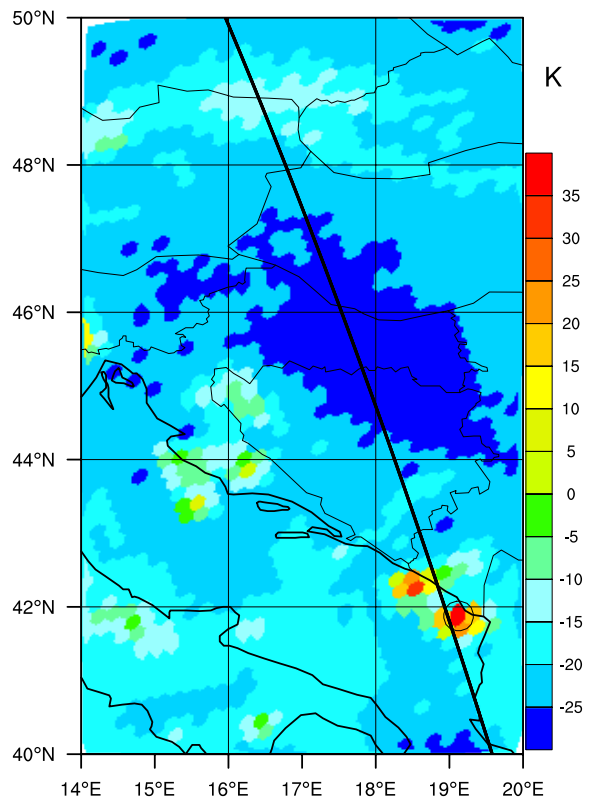


# Precipitation and AMSU-B: « Deep Convection Threshold » Analysis using TRMM

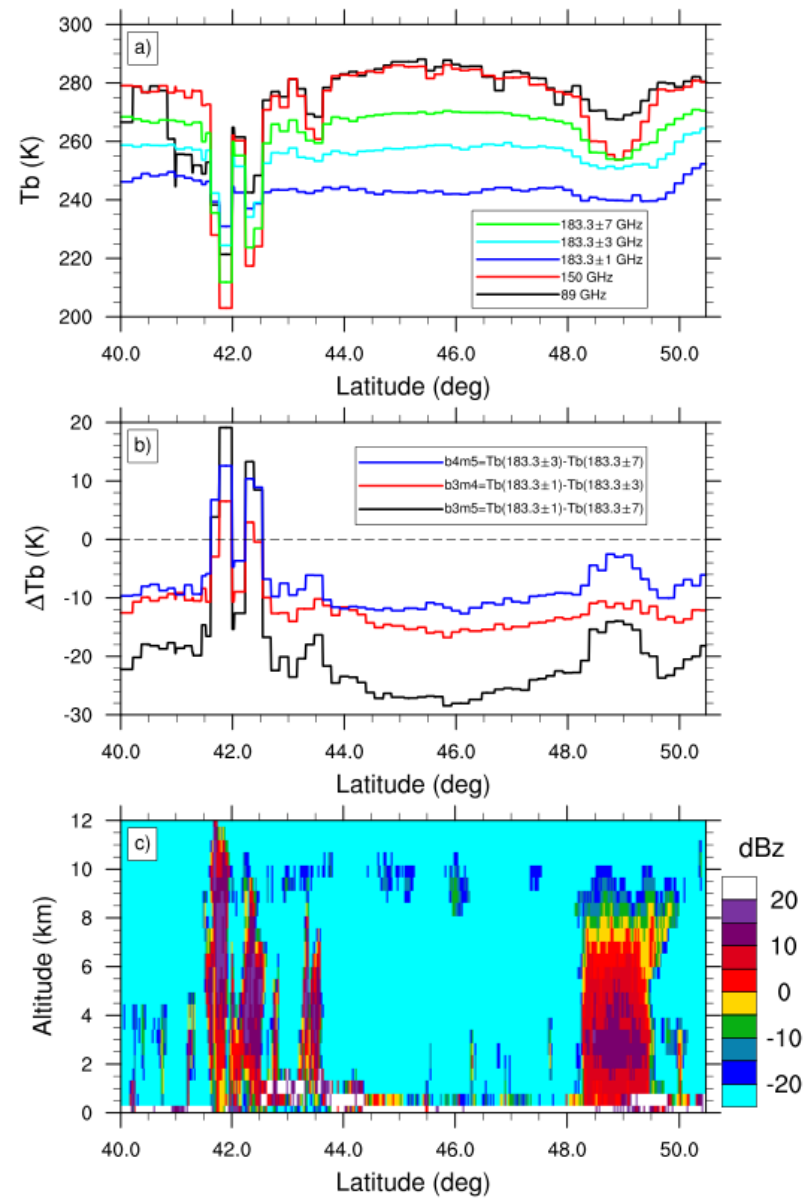
Frequency distribution for “Deep Convection Threshold” of Hong et al. (2005)

DCT = AMSU\_B45  $\geq$  0 and AMSU\_B35  $\geq$  0 and AMSU\_B34  $\geq$  0



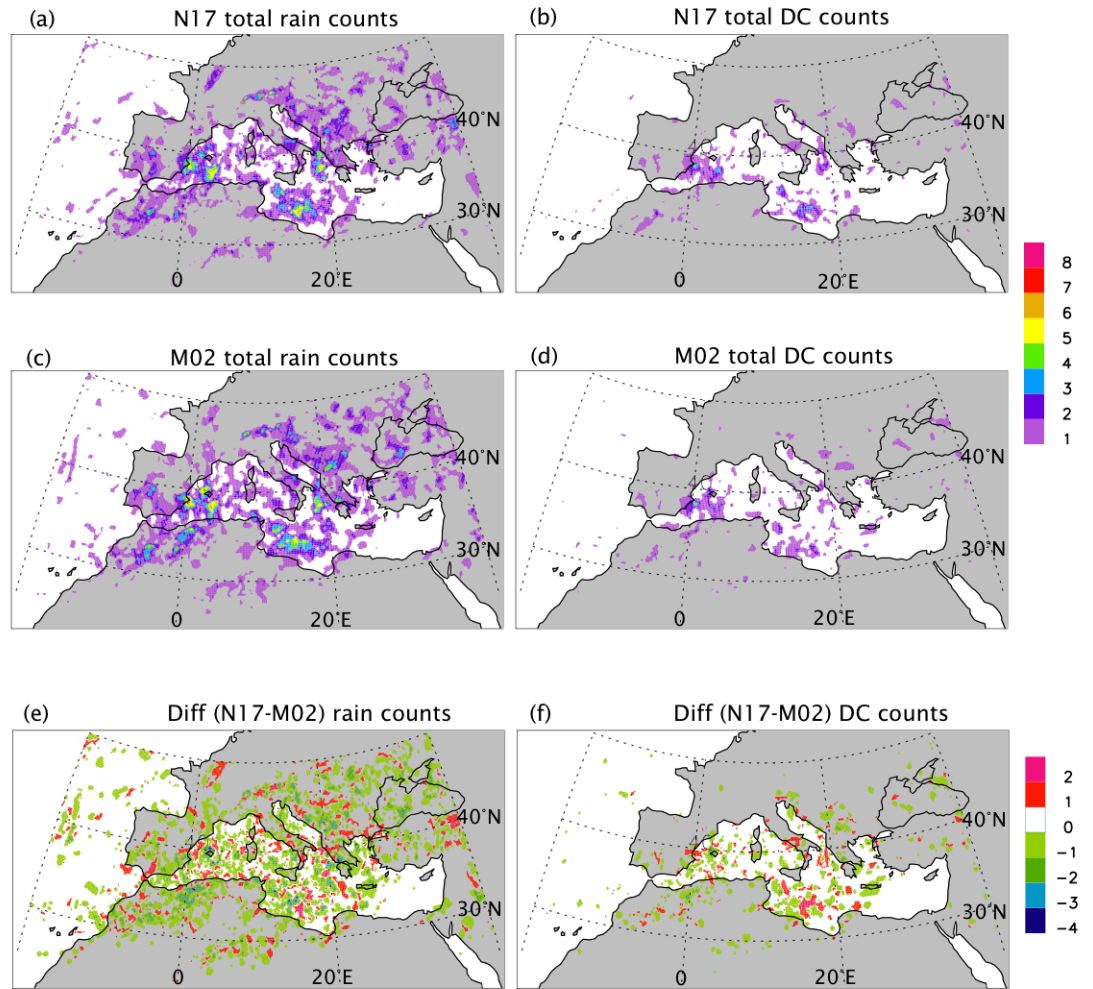


B3m5 NOAA18  
CloudSat (#10mn)



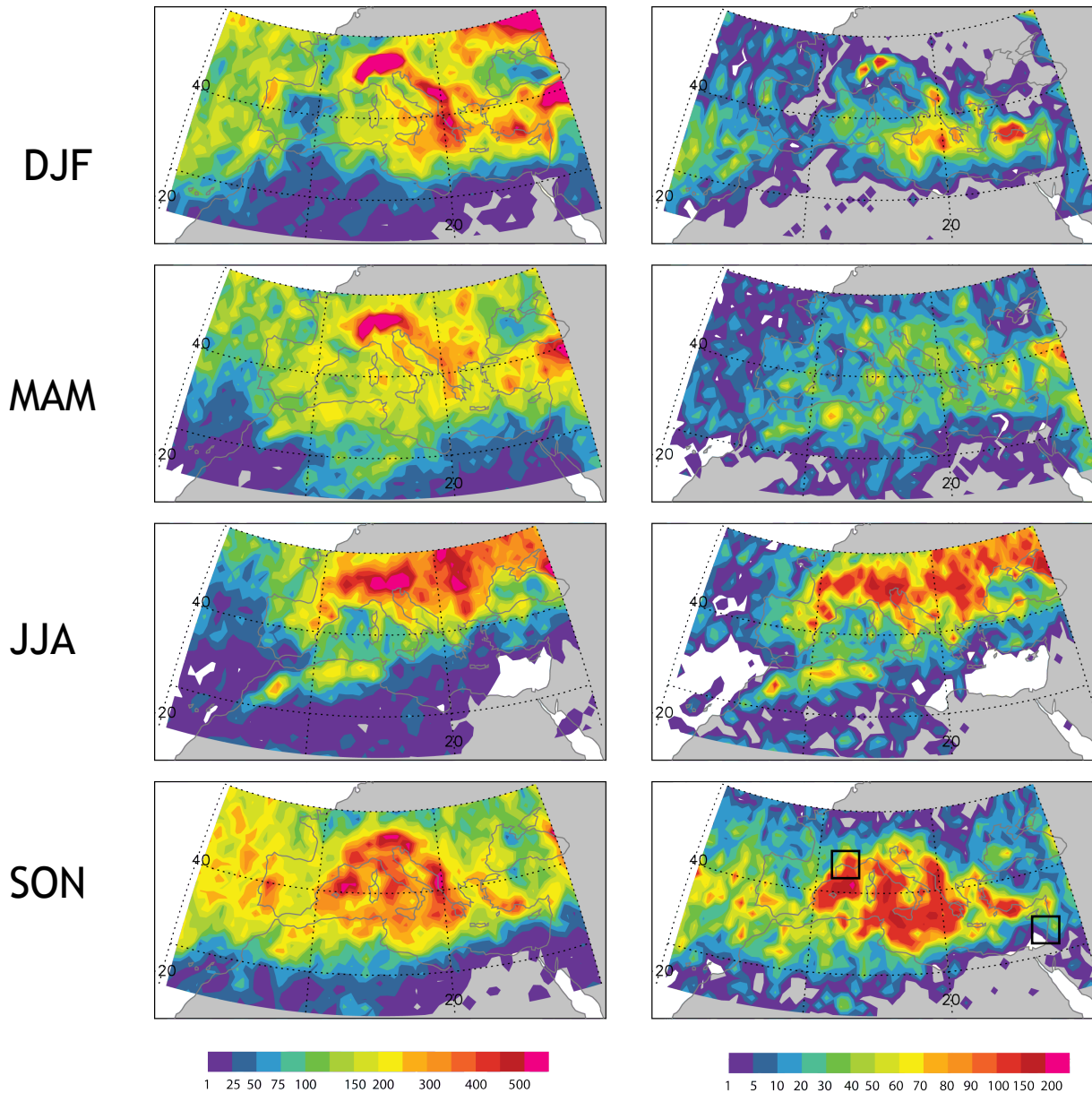
# From AMSU-B to MHS

exp: oct 2007



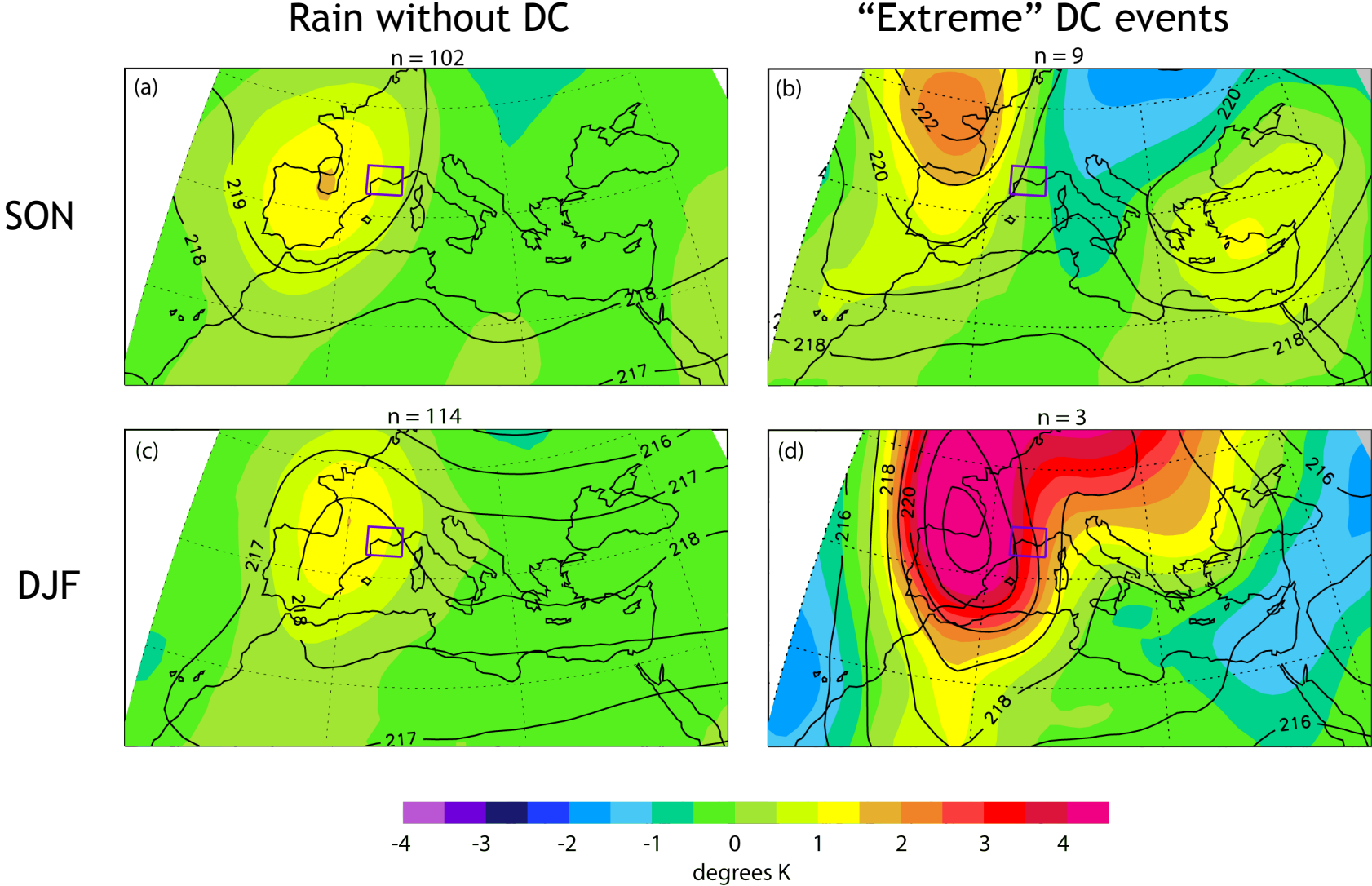
Now there is ATMS...

# Seasonal Distribution of precipitation and Deep Convection occurrences

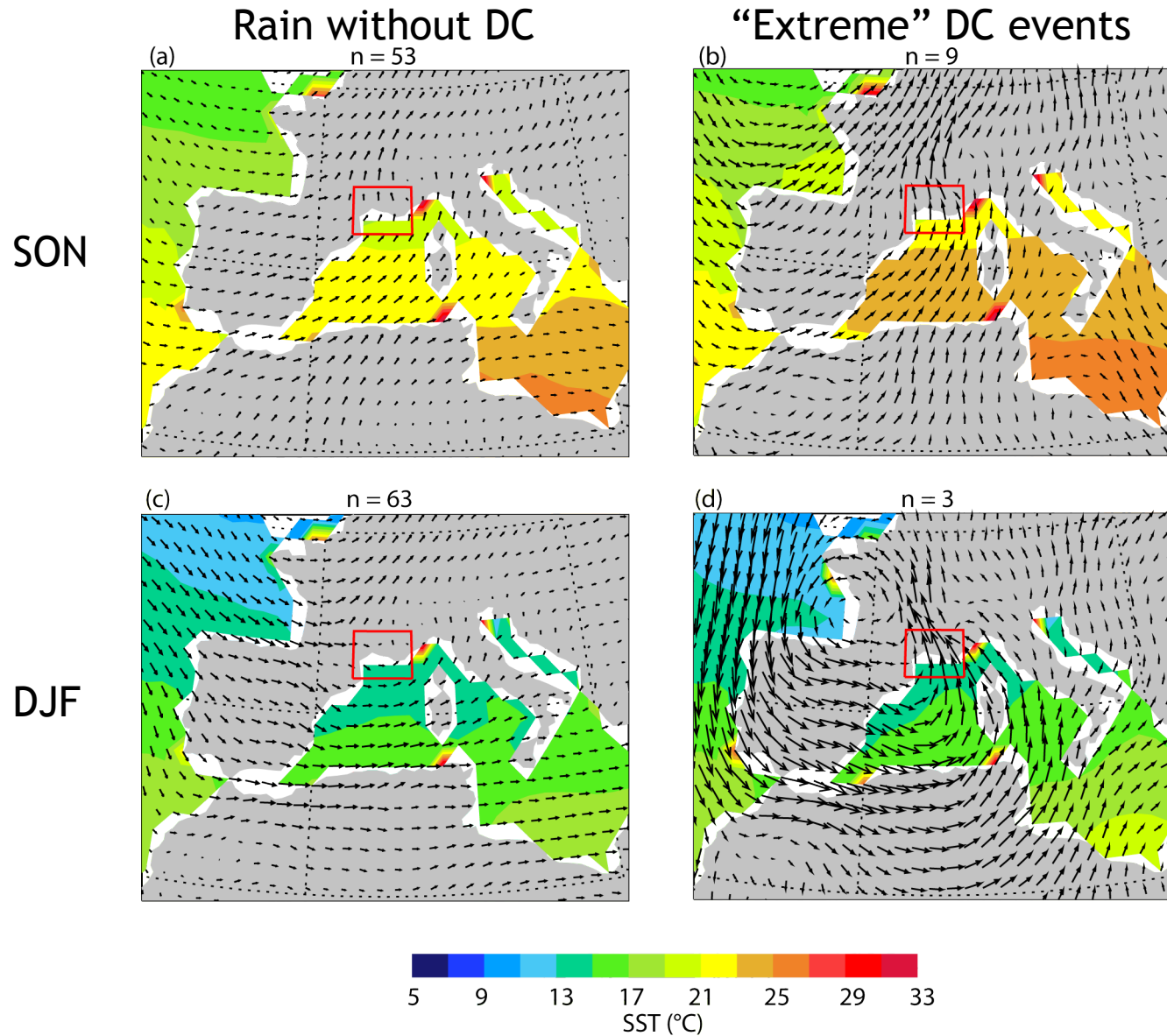


Data: NOAA-16 2001-2007

# Southern France - Upper-level structures

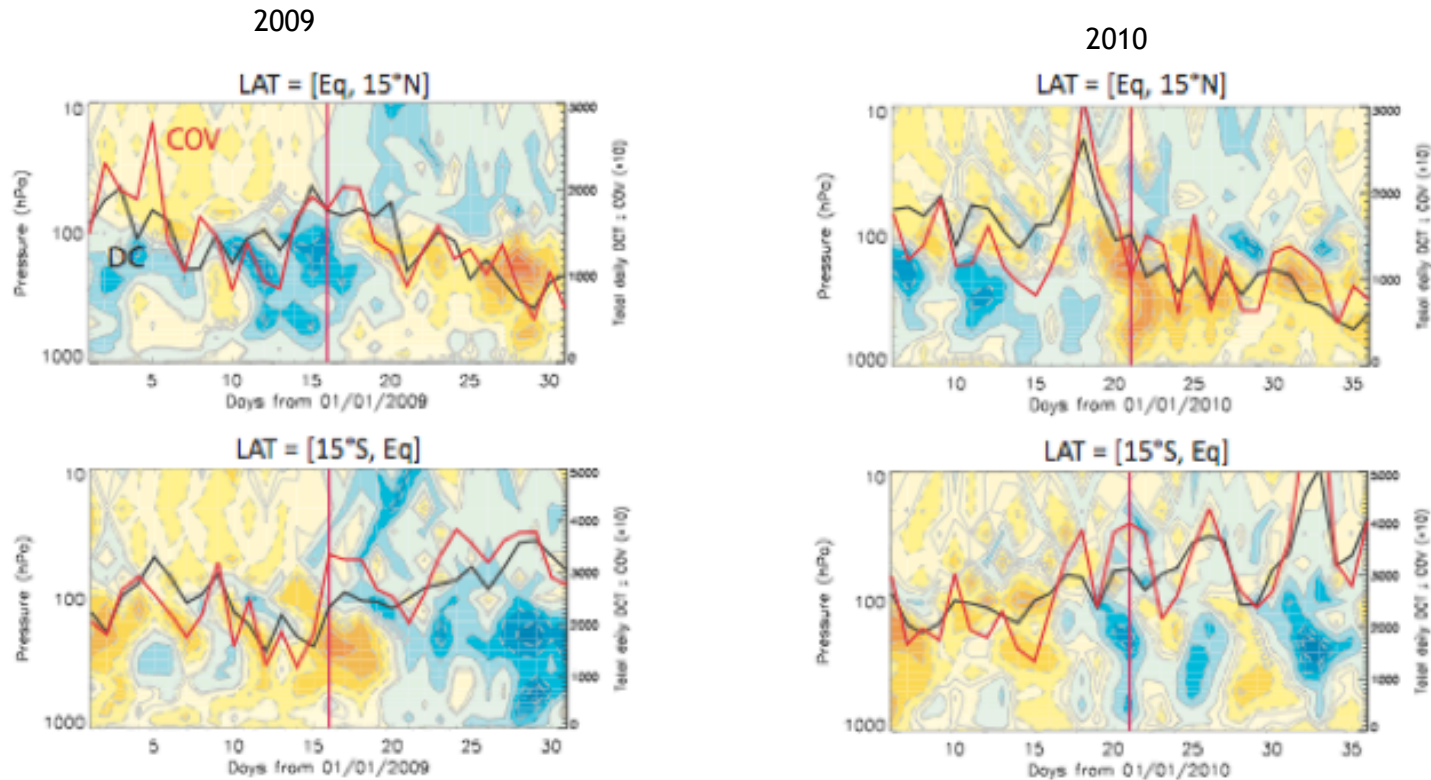


# Southern France - SST and wind at 850 hPa



SST and wind fields from Operational ECMWF analyses

## Etudes aux Tropiques : Utilisation de DC et COV pour étudier les couplages tropiques/extra-tropiques



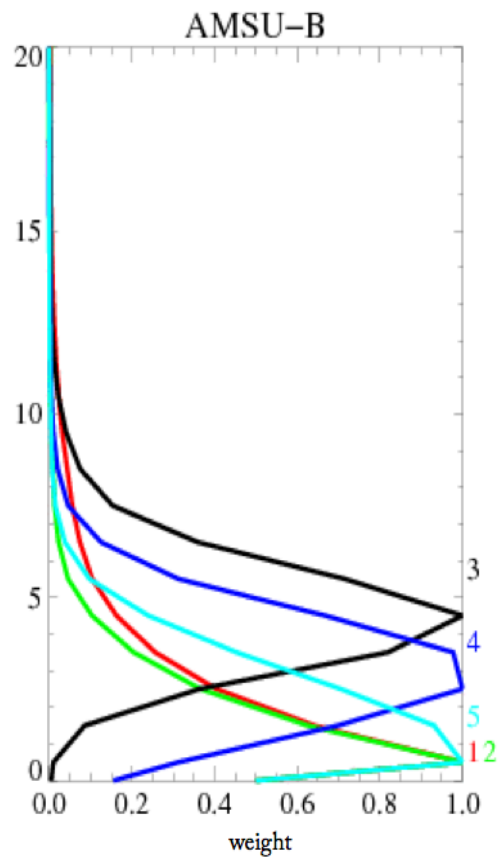
Diagrammes Hovmöller jours/altitude de l'anomalie de vitesse verticale normalisée (issue de la réanalyse Era-Interim)

-> Modulation de l'activité convective aux tropiques en lien avec les échauffements stratosphériques soudains dans la stratosphère arctique

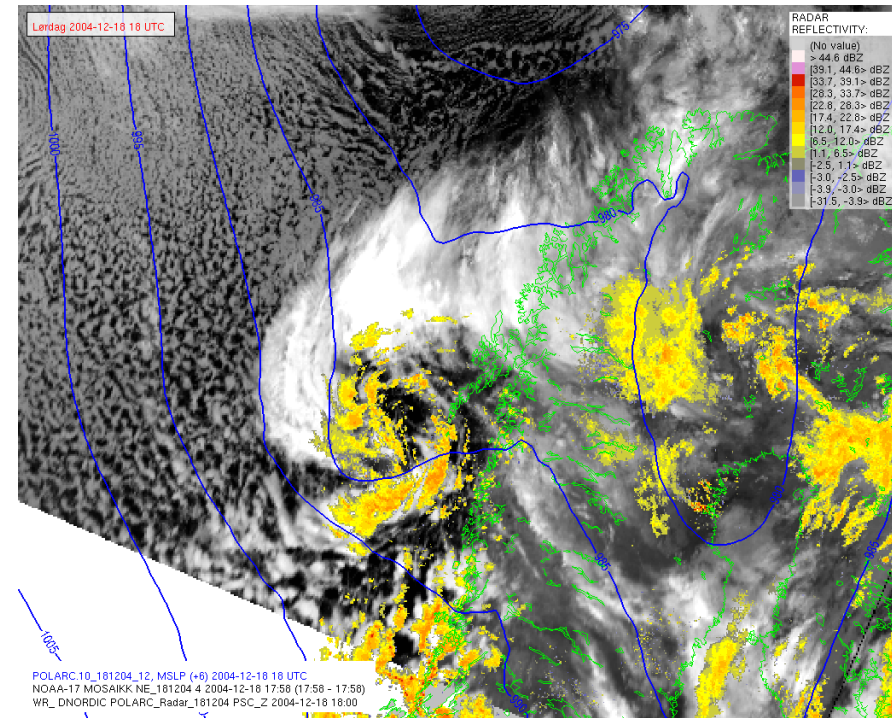
+ Etude cycle diurne convection en Amazonie: Funatsu et al, JGR, 2012

# Application à l'étude des polar lows

Cas du 12 Décembre 2008

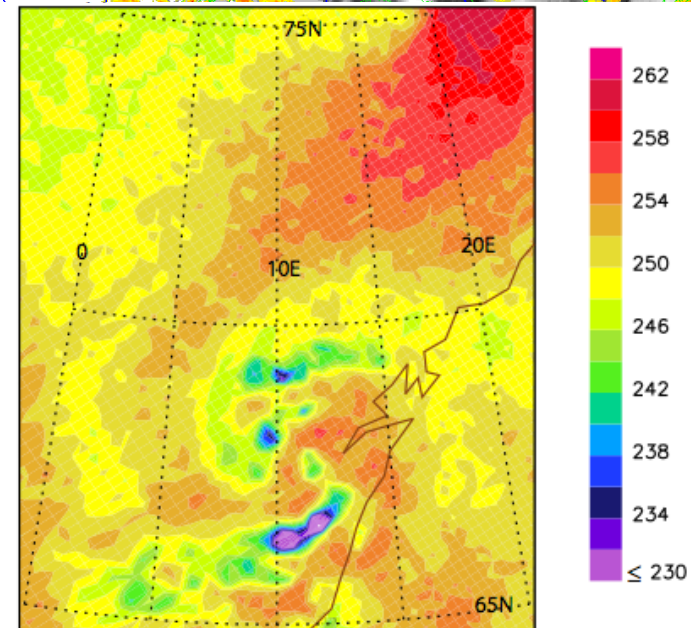


Fonctions de poids AMSU-B, profil subarctique



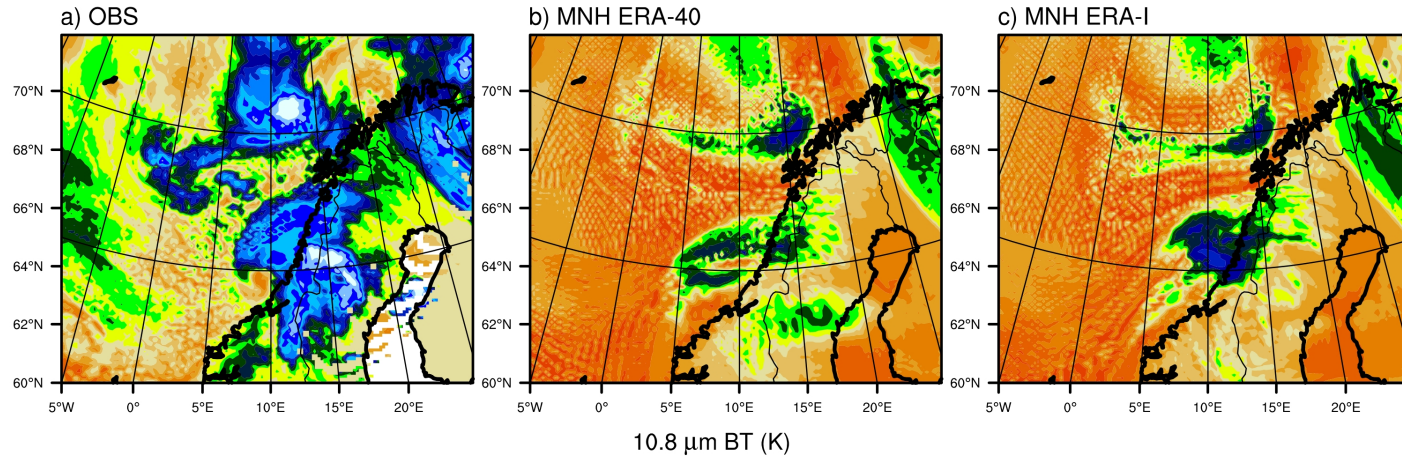
POLARC10\_181204\_12, MSLP (+6) 2004-12-18 18 UTC  
 NOAA-17 MOSAIKK NE\_181204\_4 2004-12-18 17:58 (17:58 - 17:58)  
 WR\_DNORDIC POLARC\_Radar\_181204\_PSC\_Z 2004-12-18 18:00

AMSU-B canal4

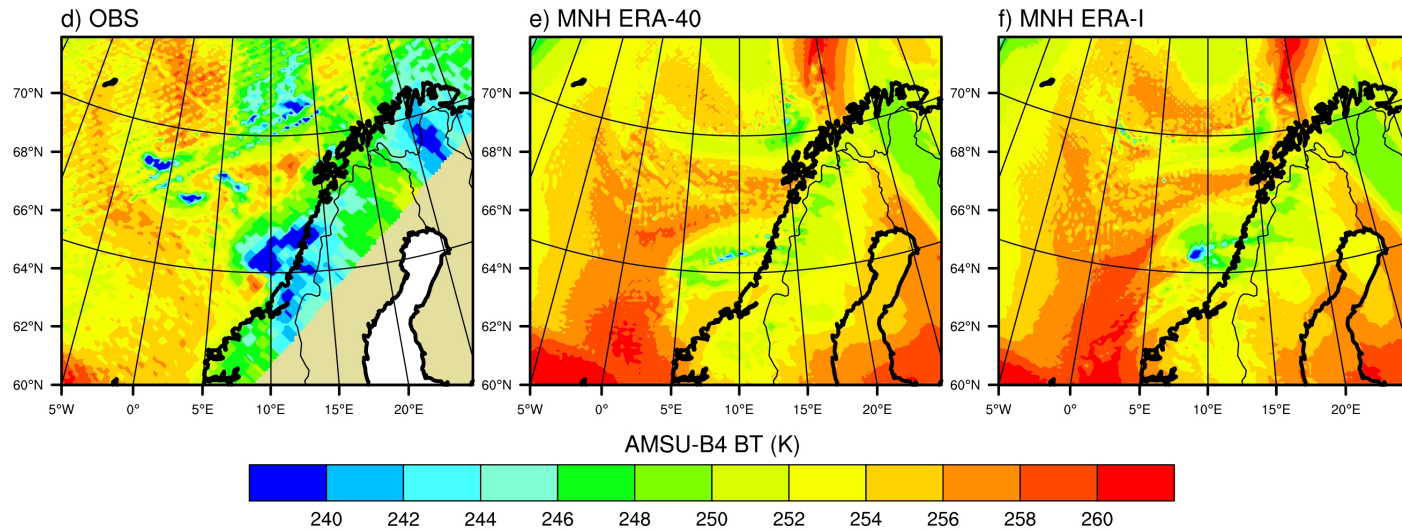




# Application à l'étude des polar lows: Validation simulations Meso-NH par une approche modèle-vers-satellite



AMSU-B4



## Conclusions & perspectives

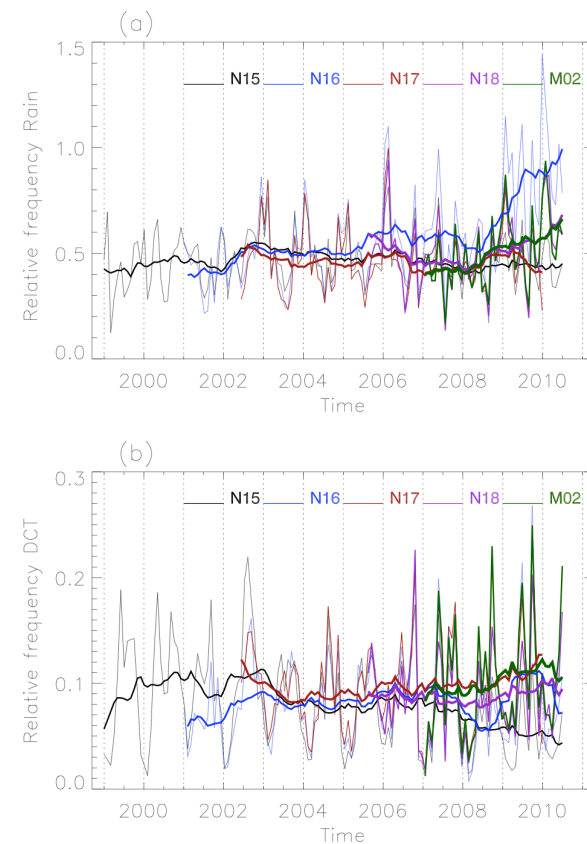
Intérêt de AMSU-B/MHS pour la détection de nuages précipitants/convectifs  
illustré pour différentes latitudes/situations

### Limitations/projets de travaux:

- Changement de radiomètres/ dérives des plateformes
- Problème de la neige (précipitation): Signal en diffusion potentiellement masqué par le signal en émission (Liu et Seo, 2013)

Solution: approche statistique couplée à CloudSat

- Problème des surfaces froides/enneigées
- ...



# Etude de la convection profonde durant la SOP HyMeX par synergie instrumentale

· Capteur IR de Meteosat  
· Réseau Euclid (détection des éclairs)

· Radar nuage aéroporté **Rasta**  
· Canaux vapeur d'eau MHS

Exemple : IOP 12a 12 octobre 2012

